

**Potential Applications of Gravity and Magnetic Methods
in Engineering and Environmental Studies**

(from Hinze, 1990, *Geotechnical and Environmental Geophysics*, v.1, p.73-126)

Target	Gravity*	Magnetics*
Mapping subsurface voids	X	x
Mapping bedrock topography	X	x
Mapping steeply dipping geologic contacts	x	X
Mapping regions of potential stress amplification (e.g. plutons and fault zone irregularities)	X	X
Mapping potential zones of weakness (e.g. paleorifts, sutures, and faults)	X	X
Mapping landfills	X	X
Mapping archeological sites	---	X
Locating buried well casings	---	X
Locating buried drums, pipelines, and other ferromagnetic objects	---	X
Locating coal burns	---	X
Determining density	X	---
Mapping groundwater cones of depression	x	---
Determining volume of organic fill in lakes	x	---
Determining groundwater volume in basins	x	---
Locating underwater ferromagnetic objects	---	X
Locating sand and gravel deposits that contain heavy minerals (including magnetite)	---	X

*X = major method; x = minor method.

Gravity: Wavelengths and Amplitude Resolution for Typical Geologic Targets

(from Yale et al., 1998, *The Leading Edge*, v.17, p.73-76)

Target	Wavelength	Amplitude
Buried cavities, tunnels, tanks	1-10 m	5-100 μ Gal
Pediment and seismic weathering layer thickness, shallow gas pockets, karst	10-200 m	0.05-0.2 mGal
Shallow salt domes and cap rock	200-1000 m	0.1-0.3 mGal
Anticlines, faults, deep salt dome flanks and overhangs	0.5-4 km	0.2-2.0 mGal
Deep sedimentary basin structure	2-20 km	5 mGal
Sedimentary basin outlines and boundaries, plate tectonic structures	10-100 km	10 mGal

Useful Equations for Gravity and Magnetic Fields
(from Hinze, 1990, *Geotechnical and Environmental Geophysics*, v.1, p.73-126)

Source	Gravity	Magnetics
Susceptibility magnetic units	---	k (SI units) = 4πk (cgs units) 1 nanoTesla (nT) = 10 ⁻⁵ Oersted
Basic	$g = \frac{G\rho Vd}{r^3}$	$T = \frac{kHV}{r^n}$, k = k _m x % magnetite
Sphere	$g = \frac{4\pi G\rho R^3}{3d^2} \frac{1}{\left[1 + (x^2/d^2)\right]^{3/2}}$	$T = \frac{8\pi kHR^3}{3d^3} \frac{\left[1 - (x^2/2d^2)\right]}{\left[1 + (x^2/d^2)\right]^{5/2}}$
Vertical Cylinder (infinite)	$g = \frac{\pi G\rho R^2}{(x^2 + d^2)^{1/2}}$	$T = \frac{\pi dkHR^2}{(x^2 + d^2)^{3/2}}$, E >> d
Horizontal Cylinder (infinite)	$g = \frac{2\pi G\rho dR^2}{(x^2 + d^2)}$	$T = \frac{2\pi kHR^2(d^2 - x^2)}{(x^2 + d^2)^2}$
Narrow Vertical Sheet	$g = \frac{2G\rho t \ln(x^2 + d_2^2)^{1/2}}{\ln(x^2 + d_1^2)^{1/2}}$	$T = 2kHt \left[\frac{d_1}{(x^2 + d_1^2)} - \frac{d_2}{(x^2 + d_2^2)} \right]$
Vertical Fault	$g = 2G\rho t\theta$	$T = \frac{2khtx}{(x^2 + d^2)}$
Finite Slab	$g = 2G\rho t(\theta_1 - \theta_2)$	$T = 2kH(\theta_1 - \theta_2)$, E >> d

Notes: g is the gravity anomaly in Gals due to a density contrast (ρ) in g/cm³ (kg/m³); T is the magnetic anomaly in Oersteds (Teslas) in a vertical polarizing magnetic field caused by a magnetic susceptibility contrast (k) in cgs units; k_m is the volume susceptibility of magnetite, 0.25 cgs units (3.14 SI); H is the vertical magnetic field in Oersteds (Teslas); G is the gravitational constant, 6.672 x 10⁻⁸ dyne-cm²/g² (6.672 x 10⁻¹¹ N-m²/kg²); V is volume; r is the distance between the observation point and the center (sphere) or axis (cylinder) or centerline; x is the horizontal distance from the observation point to the center (sphere) or axis (cylinder) or centerline (sheet) or edge (fault); d is the depth to the center (sphere) or axis (cylinder) or top (d₁) or bottom (d₂) of sheet or centerline (fault); R is radius; E is depth extent; θ is the angle from the horizontal (θ₁ to top; θ₂ to bottom). All distances in centimeters (meters). Units in parentheses are SI.